



Impact of interannual atmospheric forcing on the Mediterranean Outflow Water variability in the Atlantic Ocean

Alexandra Bozec¹, Eric Chassignet¹, Susan Lozier²

In collaboration with George Halliwell³, and Zulema Garraffo⁴

¹ Center for Ocean-Atmospheric Prediction Studies, Tallahassee, FL

² Division of Earth and Ocean Sciences, Duke University, Duke, NC

³ Atlantic Oceanographic and Meteorological Laboratory, Miami, FL

⁴ Rosenstiel School of Marine and Atmospheric Science, Miami, FL



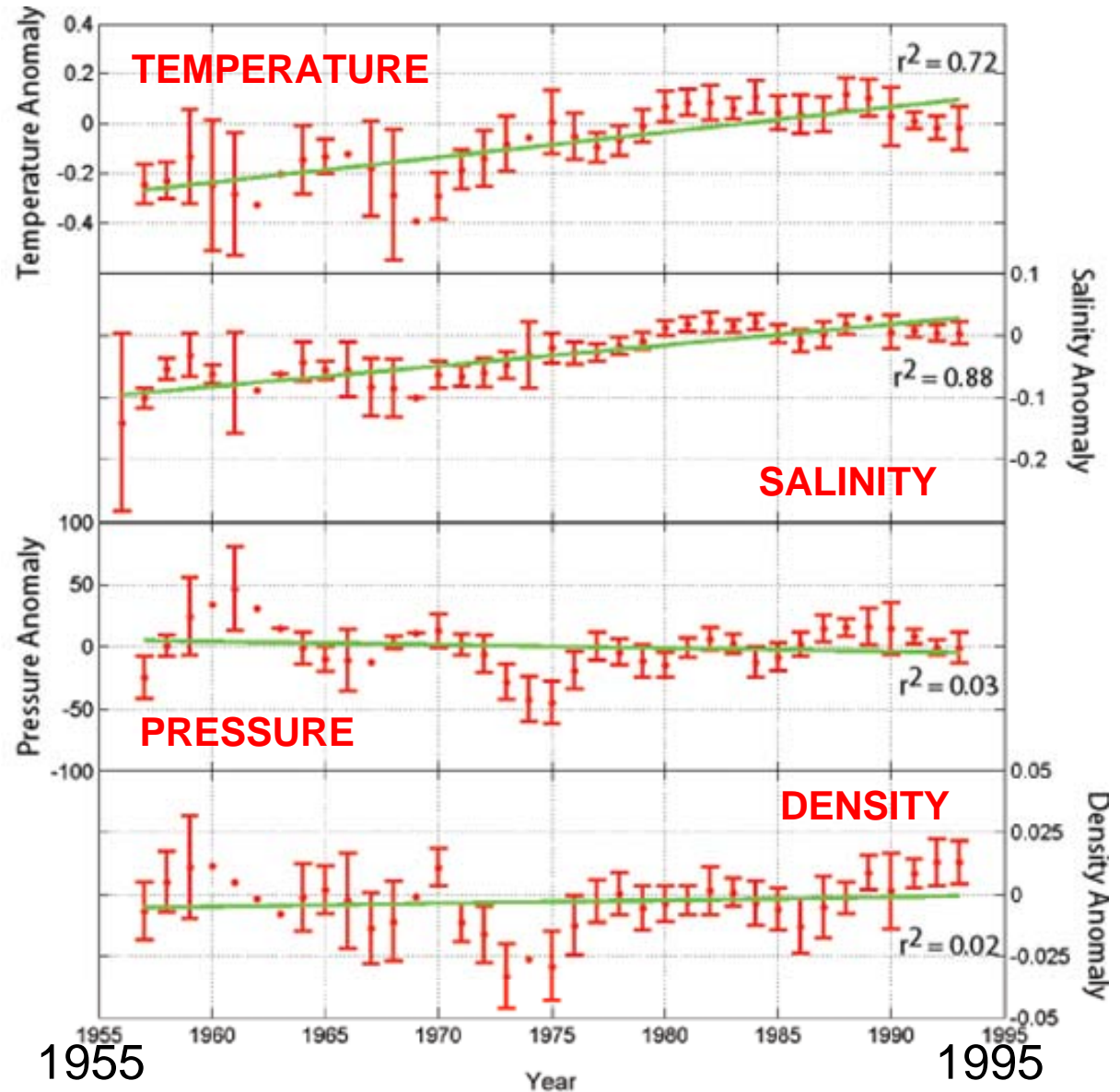
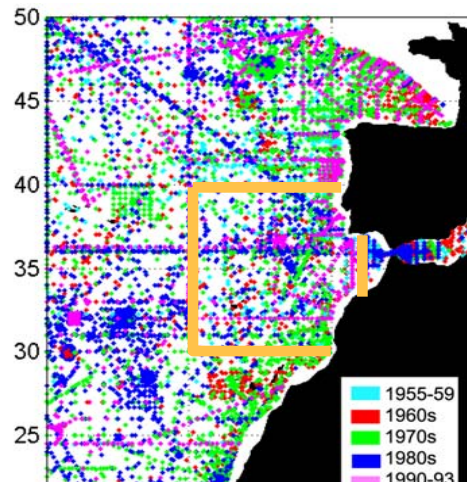
Background

Trends in water properties of the Mediterranean Outflow Water reservoir (Potter and Lozier, 2004):

Observations:

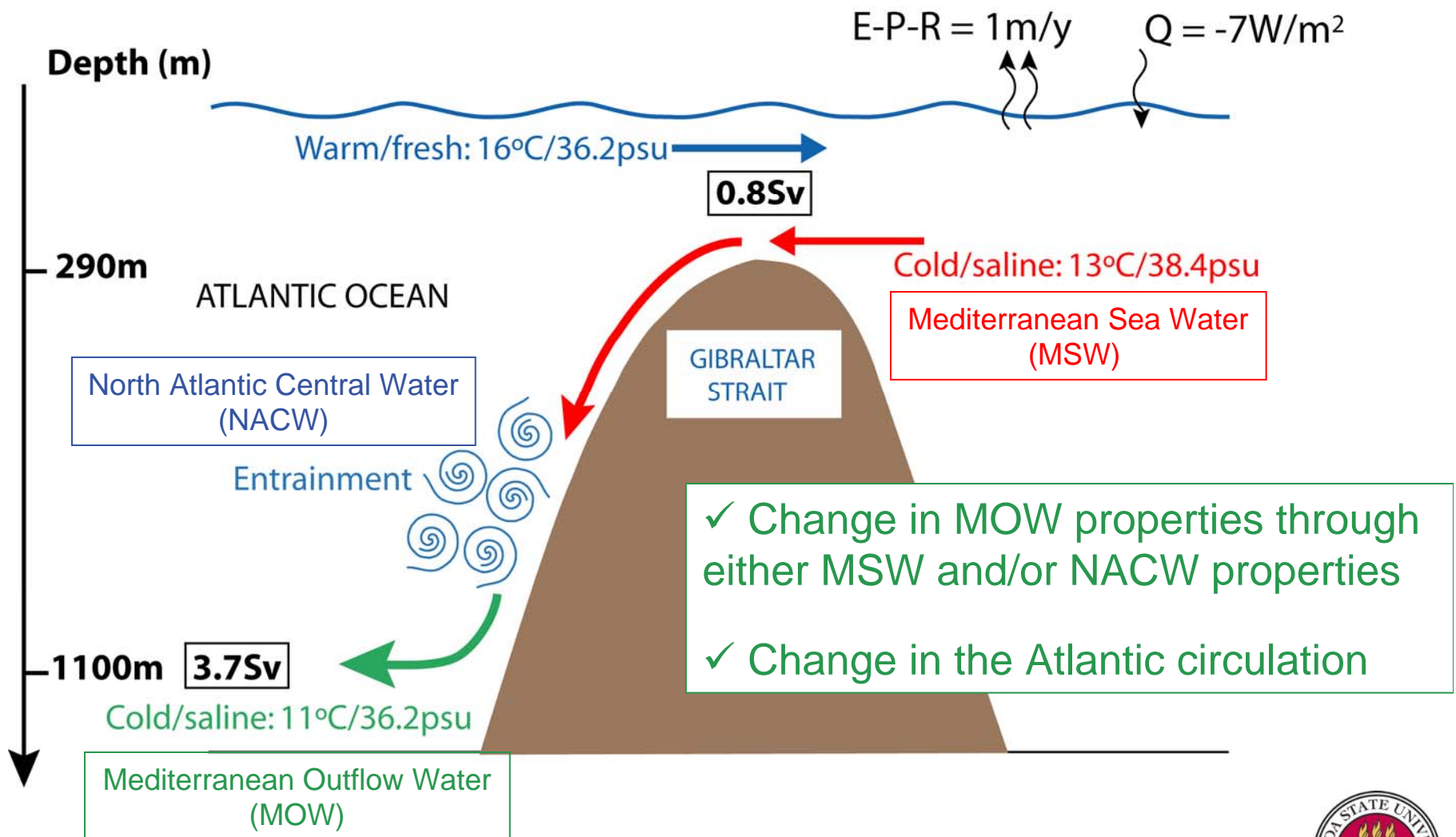
- Max of S of each profiles
- Time period : 1955 to 1993
- 3 year moving average

S-Trend (psu/decade)	T-Trend (°C/decade)
0.0283+/- 0.0067	0.101+/- 0.024





Source of Variability for the MOW Reservoir





Mediterranean Sea Water variability

Lozier and Sindlinger (2009)

Derived **MSW salinity** using different E-P products (NCEP/NCAR, ECMWF, DaSilva)

✓ NCEP/NCAR Med Sea Water:

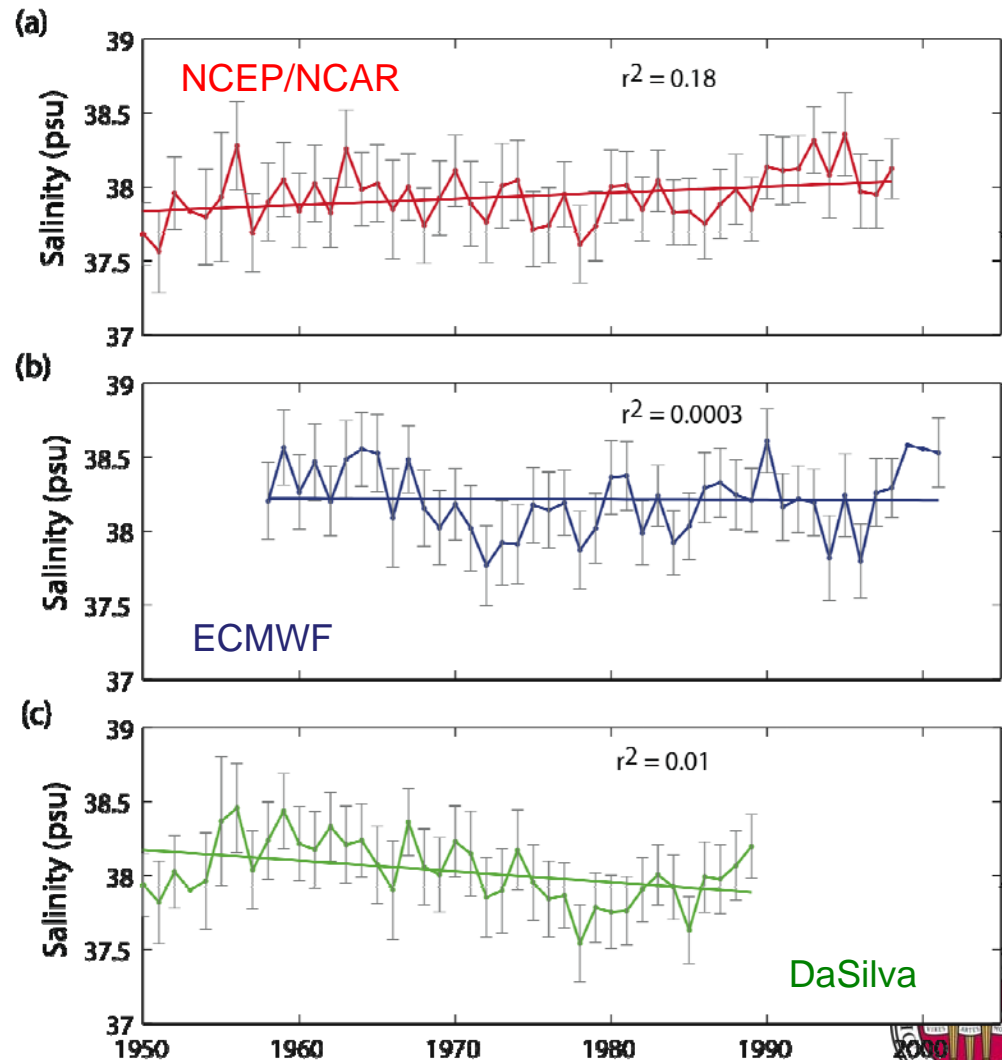
S-Trend: $+0.037 \pm 0.018$ psu/decade

✓ Resulting MOW reservoir salinity (assuming NACW constant in Price and Yang (1998)):

S-Trend: $+0.0024 \pm 0.0014$ psu/decade

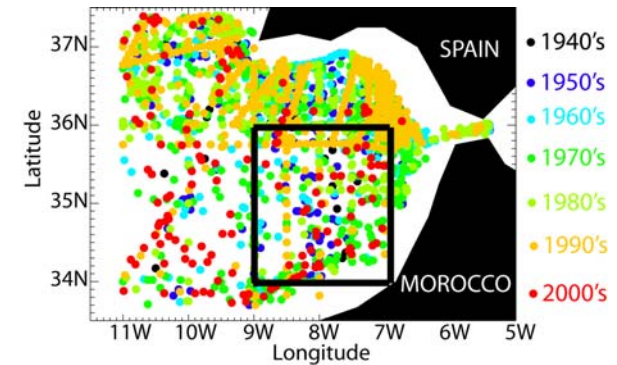
=> 10 times lower than observations

Mediterranean Sea Water from E-P

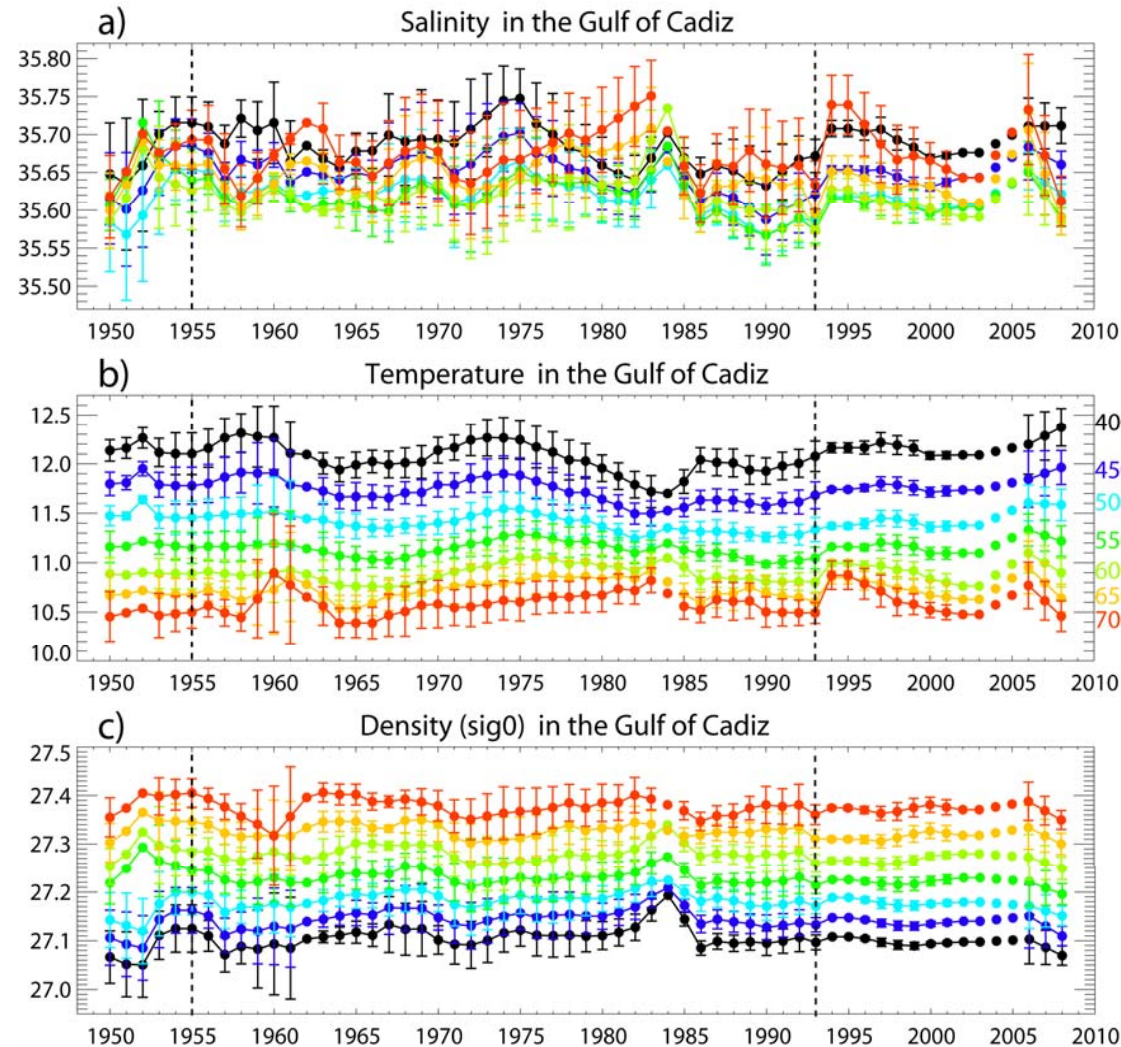




NACW entrained variability



From HydroBase 2



NACW trends between 1955-1993:

Trend S at 600m: $+0.0025 \pm 0.0090$ psu/dec.

Trend T at 600m: $+0.0069 \pm 0.0029$ °C/dec.





- ✓ The variability of the Mediterranean Sea water and of the North Atlantic Central water are too weak to be responsible for the variability of the MOW in the Atlantic.
- ✓ Hypothesis: The variability is due to the interannual variability of the atmosphere (wind and/or buoyancy forcing).





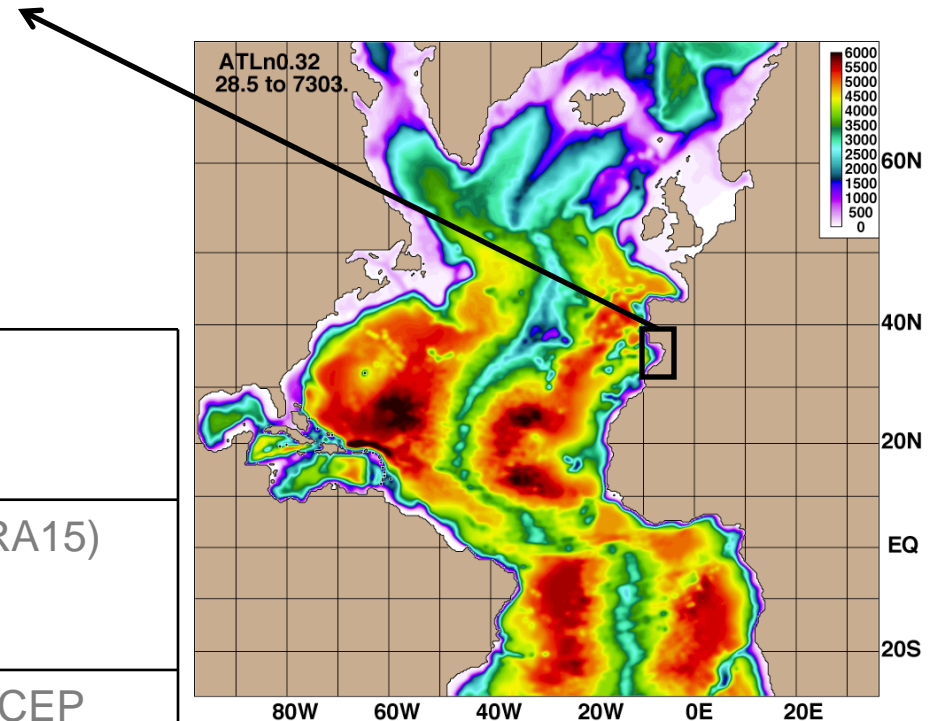
Experimental Set-up

3 simulations using $1/3^\circ$ Atlantic HYCOM with the **Marginal Sea Boundary Condition** (Price and Yang, 1998).

- ✓ Spin-up of 30 years
- ✓ 60 years simulations
- ✓ **Constant properties of the MSBC MOW:**

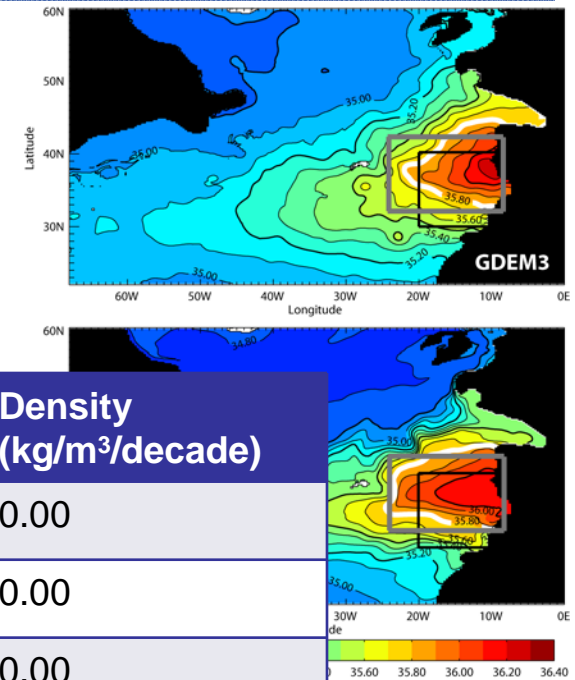
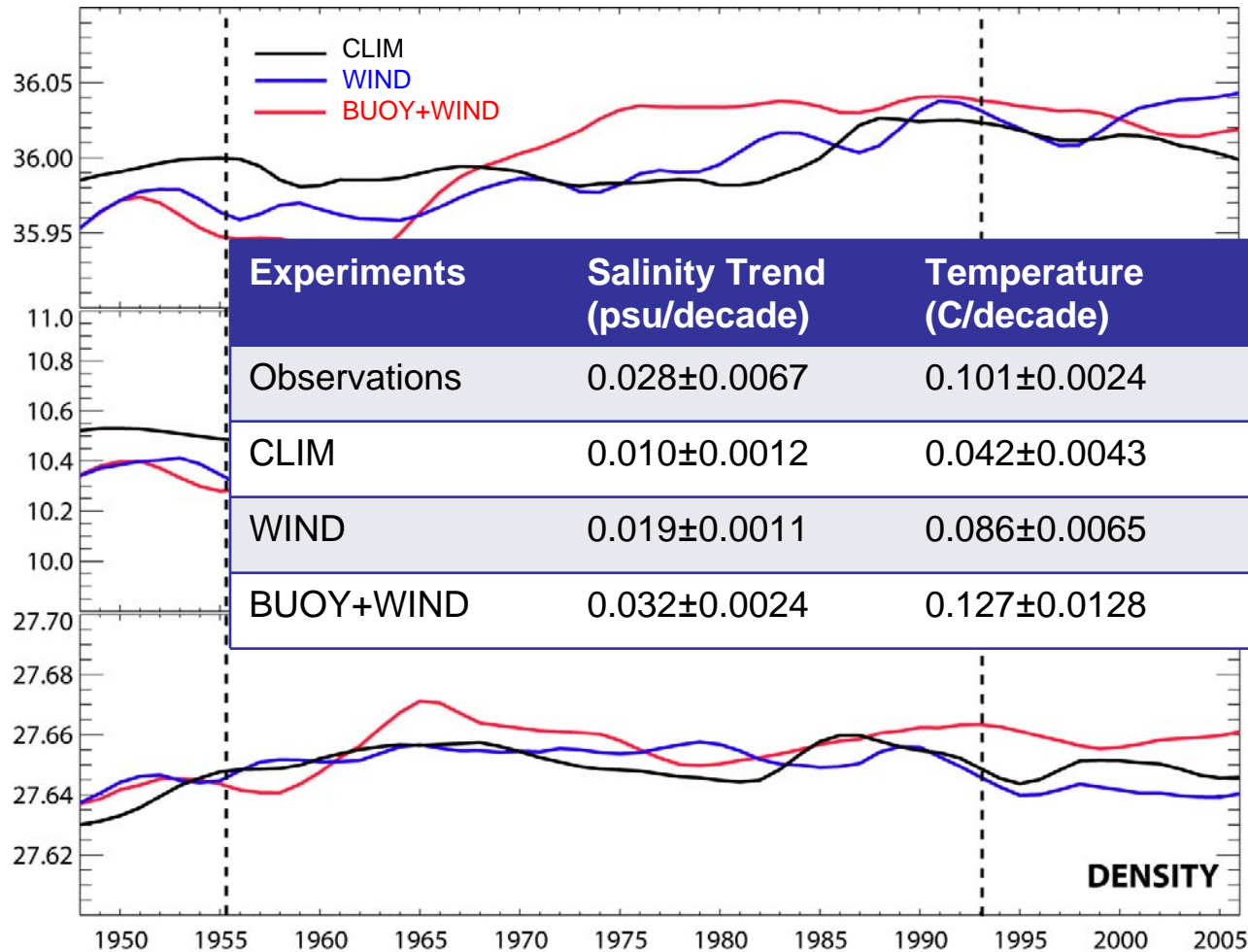
$T=11^\circ\text{C}$, $S=36.2\text{psu}$, $\text{Transport}=4\text{Sv}$.

Experiments	Atmospheric Forcing
CLIM	Climatological ECMWF (ERA15)
WIND	Interannual wind-stress (NCEP 1948-2006) and climatological ECMWF buoyancy forcing
BUOY+WIND	Interannual buoyancy and wind forcing (NCEP 1948-2006)





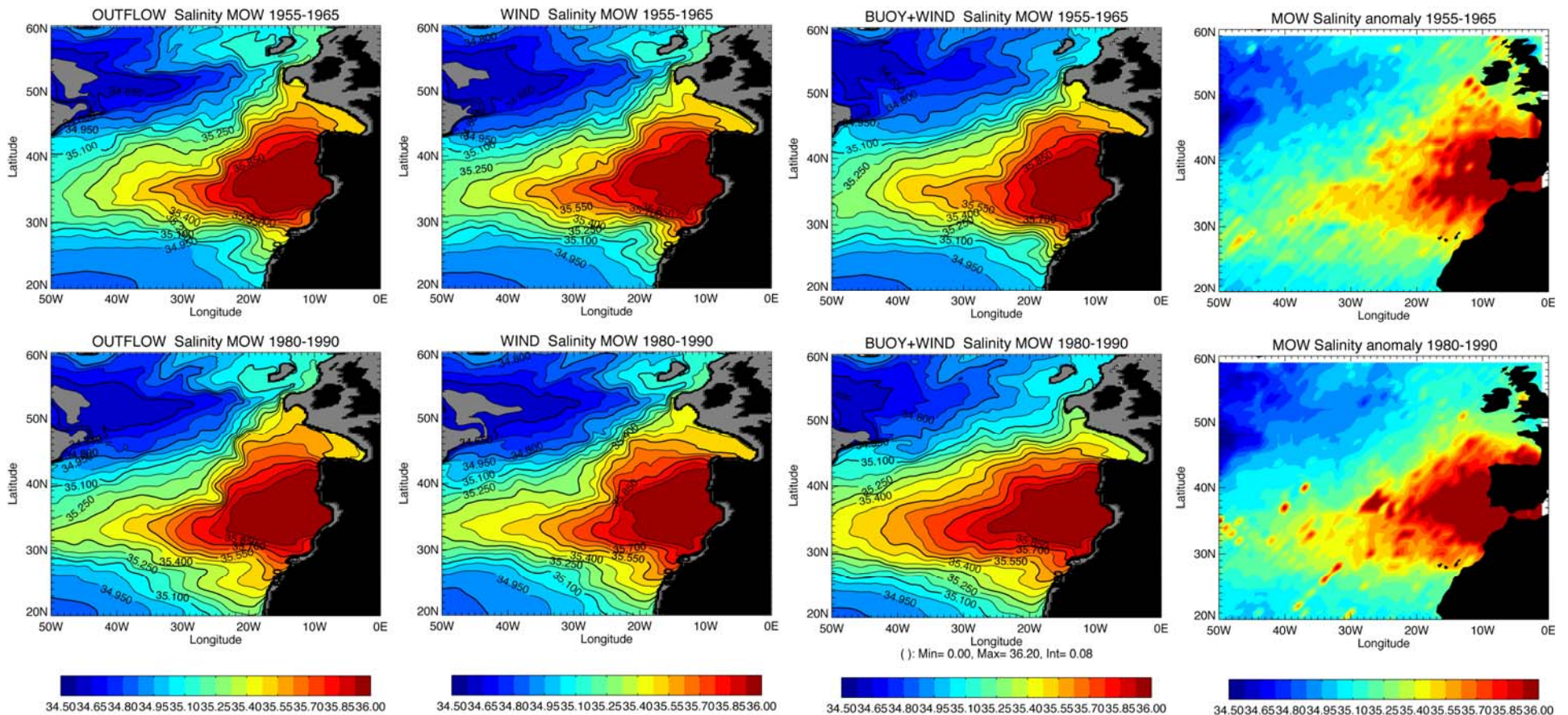
MOW Reservoir Trends in HYCOM





Spreading of the MOW

We compare the MOW tongue ($\sigma_2=36.52$) during 1955-1965 and 1980-1990 for each simulation:



CLIM

WIND

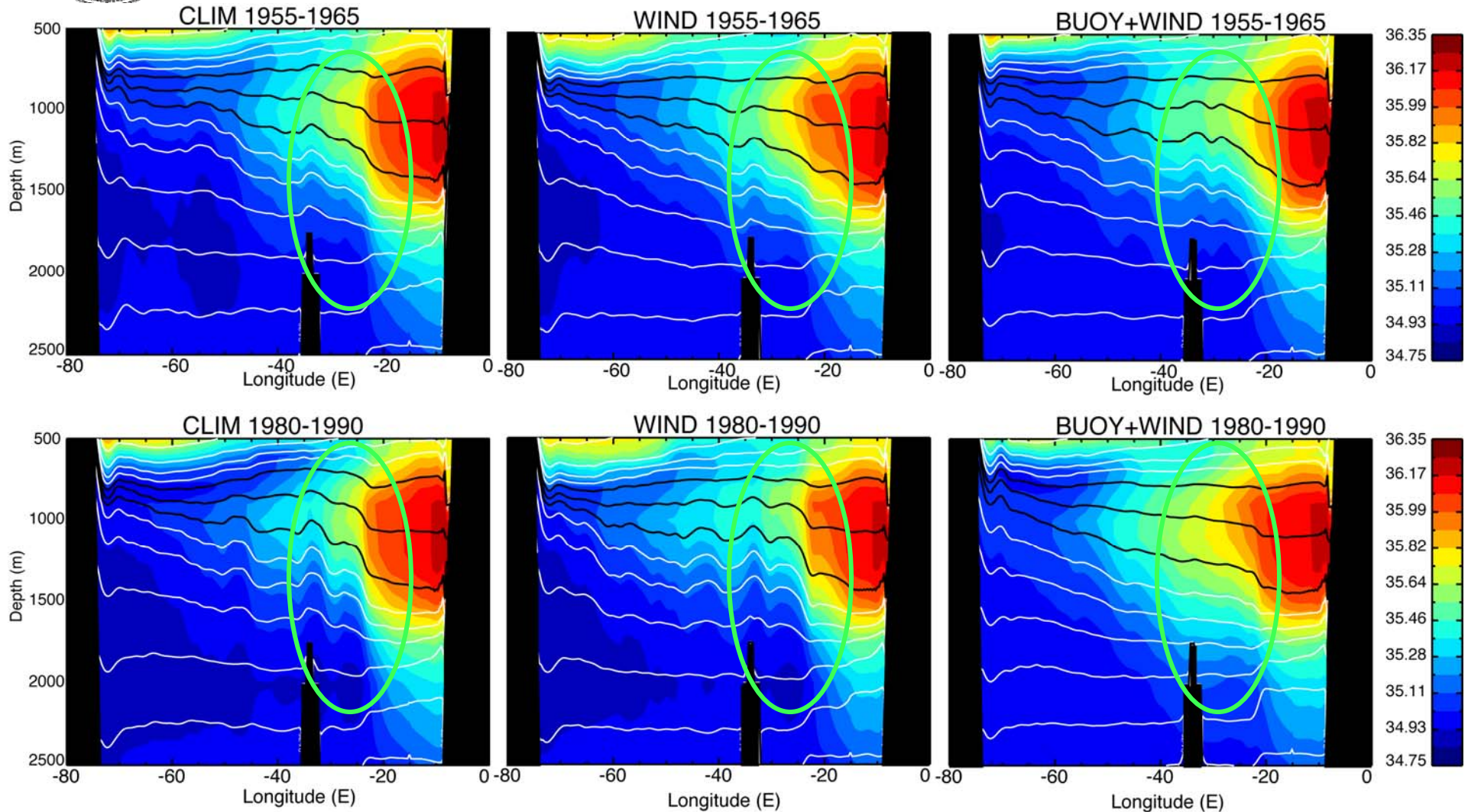
BUOY+WIND

HYDROBASE





Why is the MOW tongue expanding?



The cross section at 36°N shows tilted isopycnals in BUOY+WIND compared with the other simulations.

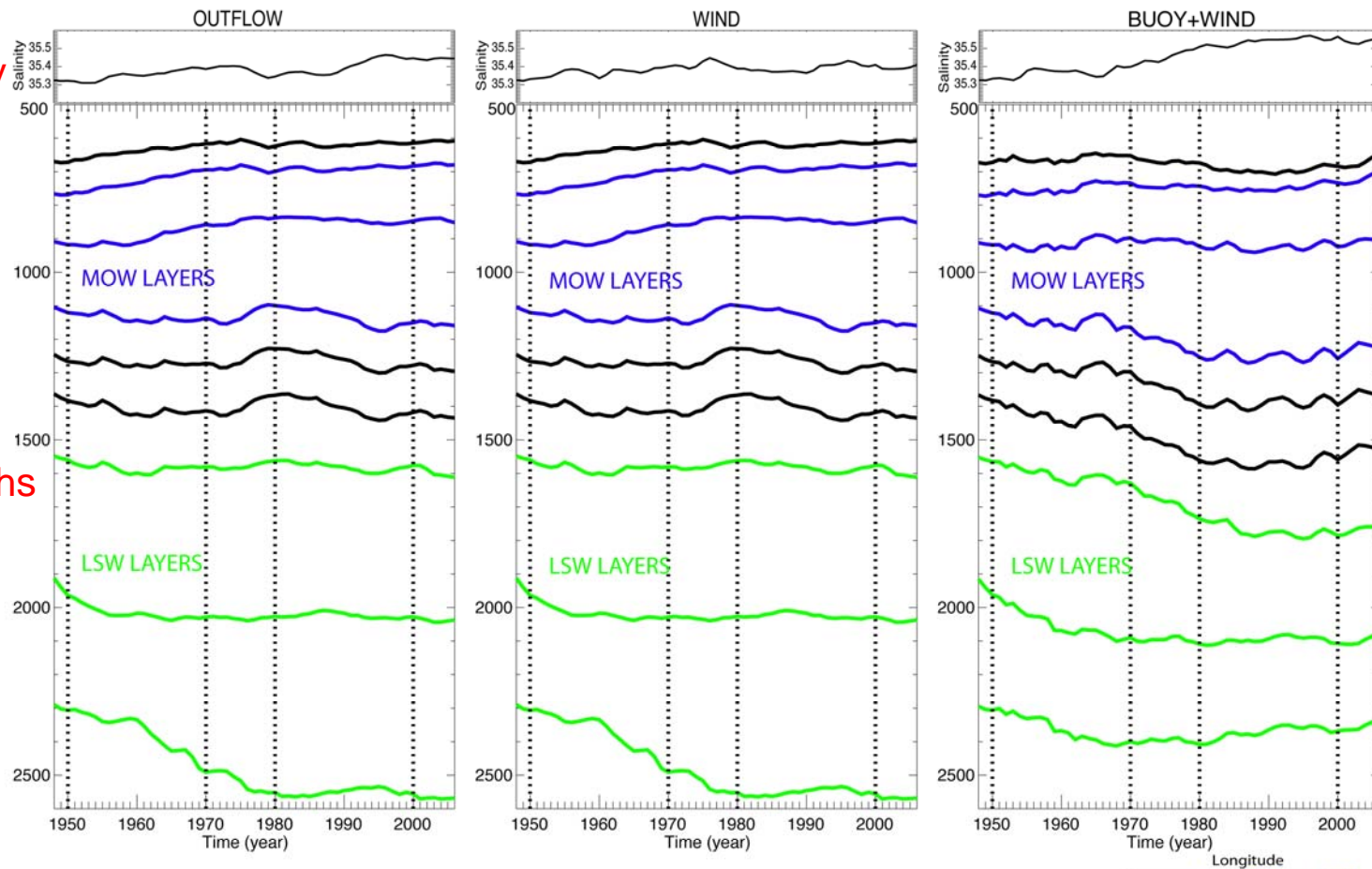




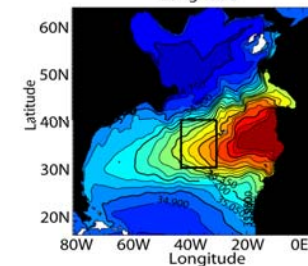
Time Evolution of Interface Depths

MOW Salinity
(psu)

Interface depths
(m)

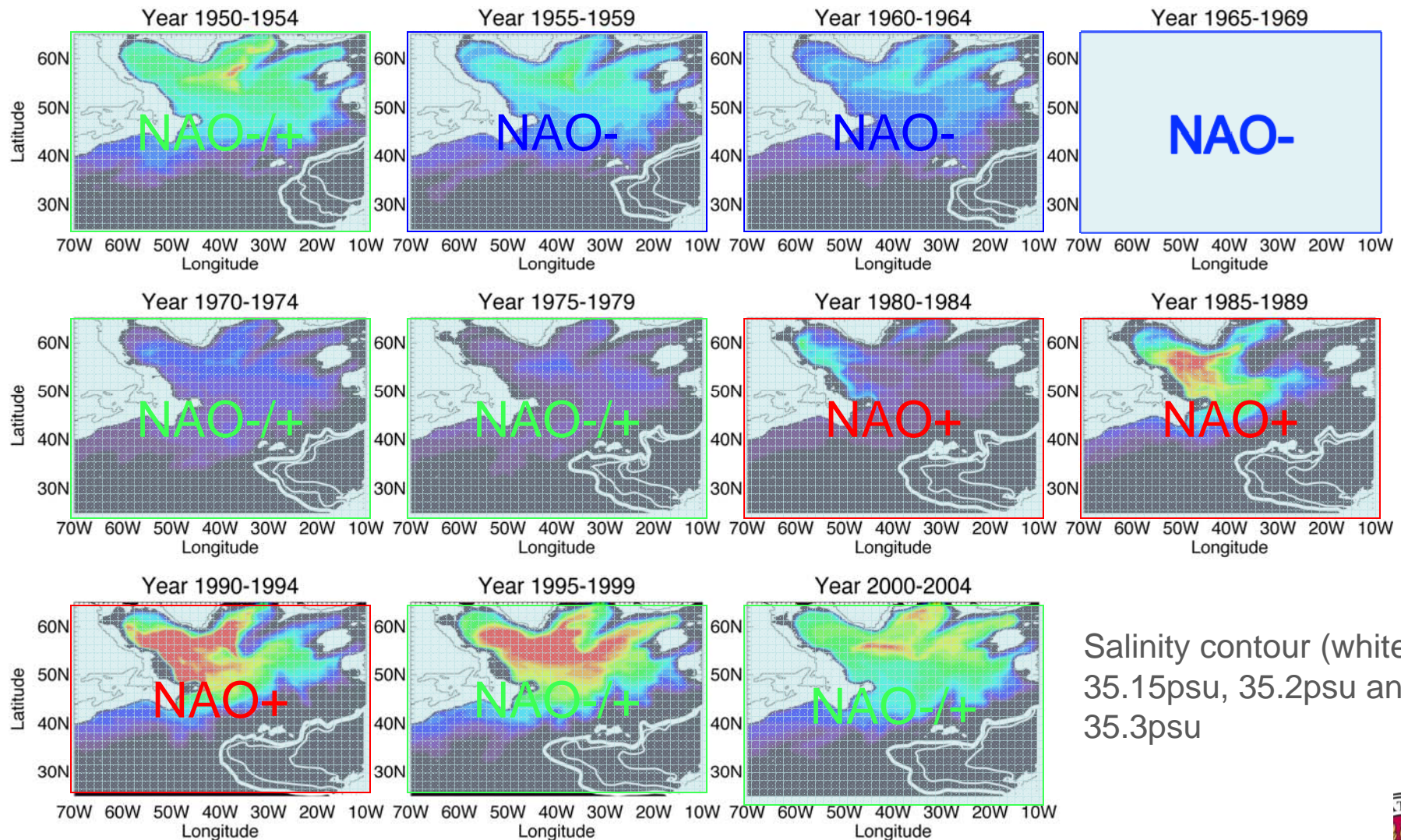


✓ MOW salinity increase coincides with a retraction of the Labrador Sea Water (LSW) from the Central Atlantic.





LSW thickness ($\sigma_2=36.83$)



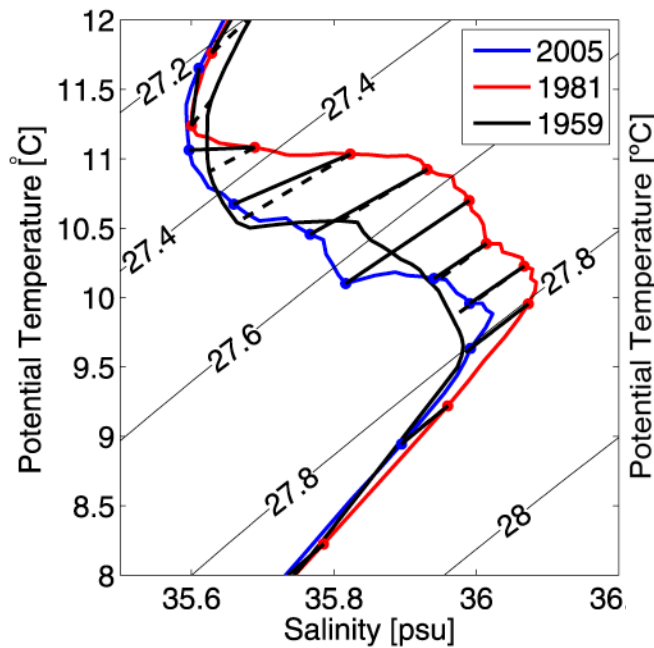
Salinity contour (white):
35.15psu, 35.2psu and
35.3psu



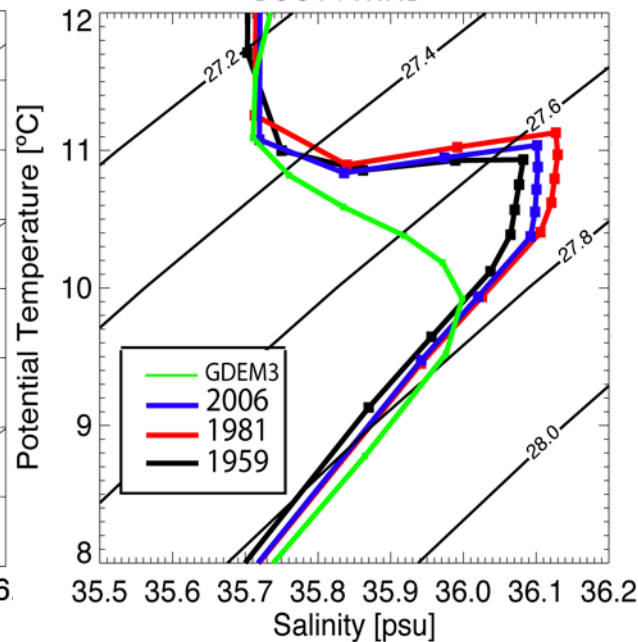


Trend reversal in the 2000s

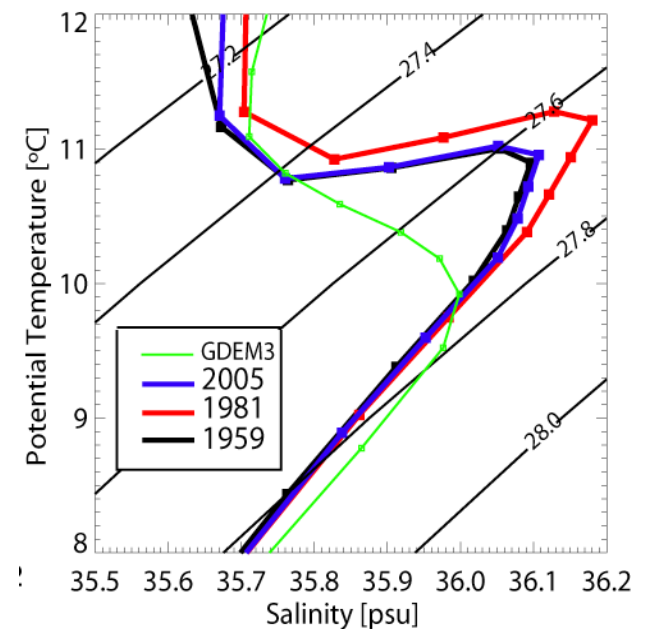
(a) Mediterranean Outflow Water ($10^{\circ} - 20^{\circ}W$)



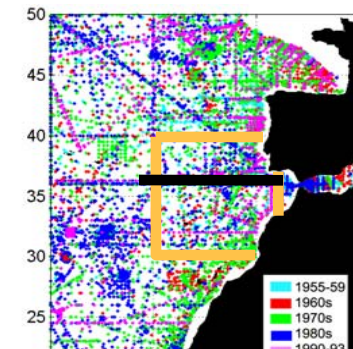
BUOY+WIND



(b) INTER experiment



From Leadbetter et al., 2007





Conclusions

- ✓ We are able to reproduce the MOW reservoir variability in the Atlantic for a constant MOW production.
- ✓ The observed MOW reservoir variability is due to circulation changes in the Atlantic Ocean induced by the atmospheric forcing.
- ✓ These circulation changes are primarily due to the variability in buoyancy forcing through the formation and flushing of Labrador Sea Water during high and low NAO periods.
- ✓ We can identify a 20-year cycle in phase with the NAO for the period 1950-2005. (i.e. time needed to fill and empty the “LSW reservoir”).

