Data assimilation in RTOFS

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Data: AVHRR, GOES and in situ - GOES bias is removed Algorithm: – Two 2DVAR analyses: one for yesterday and one for today - Time interpolated temperature is nudge while integrating from yesterday to today

 Nudge is a pseudo heat flux source/sink at the model top layer





19 Apr



Data: JASON and GFO – AVISO is not operational

 Algorithm: 2DVAR SSH, and 1D covariance in the vertical.
 SSH=MDT+SSHA with MDT from Rio (2005).



Max SHA of 0.412 at lat=37.612 lon=-71.149 and 0 other points Min SHA of -0.864 at lat=38.591 lon=-63.3 and 0 other points





Max RESID of 0.342 at lat=39.881 lon=-60.401 and 0 other points Min RESID of -0.157 at lat=39.782 lon=-60.489 and 0 other points



Results from operations as well as parallel runs give higher estimates of net transport for the section Florida to Bahamas as compared to the daily cable data. The parallel run has a similar variability to the operations but better estimates of mean transport.

Low Frequency Boundary Conditions

Internal Mode:

a) Extrapolation of velocity fluxes for advection and momentum

b) **Relaxation of Mass Fields T, S and P (interface thickness) in the buffer zones**

$$T^{k}_{t+1} = T^{k}_{t} + \Delta t \mu (\theta^{k}_{t} - T^{k}_{t})$$

$$S^{k}_{t+1} = S^{k}_{t} + \Delta t \mu (\theta^{k}_{t} - S^{k}_{t})$$

$$P^{k}_{t+1} = P^{k}_{t} + \Delta t \mu (\theta^{k}_{t} - P^{k}_{t})$$

where θ represents a slowly varying estimate (here, climatology), k is the layer and μ^{-1} is the relaxation time scale.

The width of buffer zones and values of μ^{-1} are defined a priori.

Low Frequency Boundary Conditions

External Mode:

Normal transports and elevations determined from T,S climatology and Mean Dynamic Topography using:

- a) Thermal wind relations
- b) Absolute geostrophic velocity determined by eitheri) assuming a level of no motion, or

ii) constrained by the slope of mean sea surface elevation, MDT, from Maximenko, Niiler, McWilliams (GJR,2005).

c) The mean of mean sea level is taken from MDT.

Low Frequency Boundary Conditions

Data assimilation modulates directly the low frequency mass field and sea surface height:

- a) SST: mixed layer
- b) SSH: water column and SSH

c) CTD: water column and SSH But; there is no feedback...

Large scale

 \diamond Large scales O(>500km) are well resolved by observations Internal dynamics at the mesoscale is influenced/modulated by the ambient potential vorticity Are we introducing the large scales correctly in the assimilation?



 Three simulations (from Jan 1 to March 31 2007)
 Central: No assimilation (*)
 With SST assimilation
 With SST & SSH assimilation

 Model parameters and forcing nearly identical to those used in operations







Central Jan 03-Mar 03 2007





Mean:Temperature (°C)





Gulf of Maine

46 44 42 40 -72-70-68-66-64-62

Central Jan 03-Mar 03 2007



-72-70-68-66-64-62 Diff [SST & SSH] - Central

46 44 42 40 -72-70-68-66-64-62 with SST & SSH Jan 03-Mar 03 2007





-2

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-6



Remark

 The large scale estimates derived from the model and from the combination of model and data assimilation require improvements in some geographical areas (of practical interest).

- Model
- Data
- Data assimilation