## Implicit Solution of the Vertical Diffusion Equation for the Reynolds Stress Mixing Models

The solution procedure in HYCOM follows the procedure used by Large *et al.* (1994) and is used for all three Reynolds stress vertical mixing models (KPP, GISS, MY). Decomposing model variables into mean (denoted by an overbar) and turbulent (denoted by a prime) components, the vertical diffusion equations to be solved for potential temperature, salinity, and vector momentum are

$$\frac{\partial \overline{\boldsymbol{q}}}{\partial t} = -\frac{\partial}{\partial z} \overline{w' \boldsymbol{q}'} \quad \frac{\partial \overline{S}}{\partial t} = -\frac{\partial}{\partial z} \overline{w' S'} \quad \frac{\partial \overline{\mathbf{v}}}{\partial t} = -\frac{\partial}{\partial z} \overline{w' \mathbf{v}'}.$$
(1)

Boundary layer diffusivities and viscosity parameterized as follows:

$$\overline{w'\boldsymbol{q}'} = -K_{\boldsymbol{q}} \left( \frac{\partial \overline{\boldsymbol{q}}}{\partial z} + \boldsymbol{g}_{\boldsymbol{q}} \right), \quad \overline{w'S'} = -K_{\boldsymbol{S}} \left( \frac{\partial \overline{\boldsymbol{S}}}{\partial z} + \boldsymbol{g}_{\boldsymbol{S}} \right), \quad \overline{w'v'} = -K_{\boldsymbol{m}} \left( \frac{\partial \overline{\boldsymbol{v}}}{\partial z} + \boldsymbol{g}_{\boldsymbol{m}} \right), \quad (2)$$

where the g terms represent nonlocal fluxes. For example, the KPP model includes nonlocal terms for q and S, but not for momentum. The following solution procedure is valid for any mixing model in HYCOM that calculates the diffusivity/viscosity profiles at model interfaces, whether or not nonlocal terms are parameterized.

The following matrix problems are formulated and solved:

$$\mathbf{A}_{\mathbf{T}} \Theta^{t+1} = \Theta^{t} + \mathbf{H}_{\Theta} \quad \mathbf{A}_{\mathbf{S}} \mathbf{S}^{t+1} = \mathbf{S}^{t} + \mathbf{H}_{\mathbf{S}} \quad \mathbf{A}_{\mathbf{M}} \mathbf{M}^{t+1} = \mathbf{M}^{t} + \mathbf{H}_{\mathbf{M}},$$
(3)

where superscripts t, t+1 denote model times, and **M** is the vector of a momentum component, either u or v. The matrices **A** are tri-diagonal coefficient matrices, while the vectors  $\mathbf{H}_{T}$  and  $\mathbf{H}_{s}$  represent the nonlocal flux terms. Given K model layers with nonzero thickness, where an individual layer k of thickness  $dp_{k}$  is bounded above and below by interfaces located at pressures  $p_{k}$  and  $p_{k+1}$ , the matrix  $\mathbf{A}_{s}$  is determined as follows:

$$\mathbf{A}_{\mathbf{S}}^{1,1} = \left(1 + \Omega_{S1}^{+}\right)$$

$$\mathbf{A}_{\mathbf{S}}^{k,k-1} = -\Omega_{Sk}^{-} \qquad 2 \le k \le \mathbf{K}$$

$$\mathbf{A}_{\mathbf{S}}^{k,k} = \left(1 + \Omega_{Sk}^{-} + \Omega_{Sk}^{+}\right) \qquad 2 \le k \le \mathbf{K} \quad ,$$

$$\mathbf{A}_{\mathbf{s}}^{k,k+1} = -\Omega_{Sk}^{+} \qquad 1 \le k \le \mathbf{K} - 1$$

$$(4)$$

with

$$\Omega_{Sk}^{-} = \frac{\Delta t}{dp_{k}} \frac{K_{S}(p_{k})}{(p_{k+0.5} - p_{k-0.5})},$$

$$\Omega_{Sk}^{+} = \frac{\Delta t}{dp_{k}} \frac{K_{S}(p_{k+1})}{(p_{k+1.5} - p_{k+0.5})},$$
(5)

where  $p_{k+0.5}$  represents the central pressure depth of model layer k. The nonlocal flux arrays are calculated using

$$H_{S1} = \frac{\Delta t}{\boldsymbol{d}p_1} K_S(p_{k+1}) \boldsymbol{g}_S(p_{k+1})$$

$$H_{Sk} = \frac{\Delta t}{\boldsymbol{d}p_k} \Big[ K_S(p_{k+1}) \boldsymbol{g}_S(p_{k+1}) - K_S(p_k) \boldsymbol{g}_S(p_k) \Big] \quad 2 \le k \le \mathrm{K} \,.$$
(6)

The solution is then found by inverting the tri-diagonal matrix  $\mathbf{A}$ . The matrix problems are formulated and solved in the same manner for potential temperature and momentum components.

## REFERENCE

Large, W. G., J. C. Mc Williams, and S. C. Doney, 1994: Oceanic vertical mixing: a review and a model with a nonlocal boundary layer parameterization. *Rev. Geophys.* **32**, 363-403.